

SURFICIAL GEOLOGY OF HODGEMAN COUNTY, KANSAS

Geology by William C. Johnson and Terri L. Woodburn 2012

Funded in part by the USGS National Cooperative Geologic Mapping Program

GENERAL GEOLOGY

Hodgeman County, in southwest Kansas, is bounded by Ness County to the north, Pawnee and Edwards counties to the east, Ford County to the south, and Finney and Gray counties to the west. The county has a total area of 864 mi² (2,238 km²). Cretaceous (145-65 million years ago) and Neogene (23-2.6 million years ago) sedimentary rocks—sandstones, limestones, shales, cherts, and conglomerates—crop out in the county and range in age from lower Cretaceous (Dakota Formation) to the Miocene and early Pliocene (Ogallala Formation) (Ludvigson et al., 2009). Pleistocene and Holocene loess (wind-deposited silt) mantles most of the southern and eastern portions of the upland regions of the county.

GEOMORPHOLOGY

Hodgeman County lies within the High Plains and Smoky Hills physiographic regions (Schowee, 1949). In this county, the High Plains are characterized by relatively flat topography and consist of the loess and Ogallala-capped uplands (Moss, 1932; Schowee, 1949). In contrast, the Smoky Hills consist primarily of Cretaceous and Greenhorn Limestone uplands and intermediate slopes (Moss, 1932; Schowee, 1949). Topographic relief in the county is 594.8 ft (183 m), with the highest area (2,677 ft, 816 m) in the northwest corner of the county near the Fort Hays limestone exposure, and the lowest point (2,083 feet, 635 m) where the Pawnee River leaves Hodgeman County in the northeast corner. The primary drainage, the Pawnee River, flows from west to east along the northern boundary of the county with its principal tributary, Backster Creek, flowing southward to it through the southwest and central portions of the county. The south and south-central sections of the county are drained by Saw Log Creek, a tributary to Backster Creek. These major drainages have become entrenched, creating broad terraces with narrow floodplains. Three lakes are located in the central part of the county—Hance that occupies the central part of Backster Creek west of Jetermo, Jetermo City Lake, on Spring Creek southwest of Jetermo; and Hodgeman County State Lake, on Rock Creek southwest of Jetermo.

MINERAL RESOURCES

Mineral resources in Hodgeman County include sand, gravel, and calcite from the Ogallala Formation, and sand and gravel from alluvial deposits, building stone quarried from the Greenhorn Limestone, and oil and gas. Sand and gravel deposits are used chiefly for road surfacing. Utilization of the Forcepoint limestone bed of the Greenhorn was much greater in the past when it was exploited for structural work such as buildings and bridges, but most frequently for fence posts (Moss, 1932).

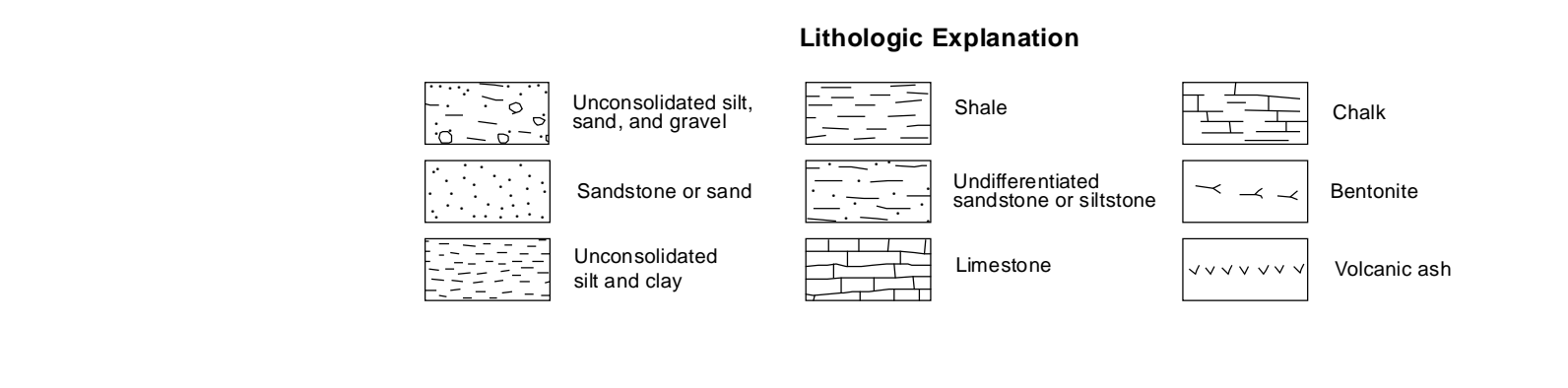
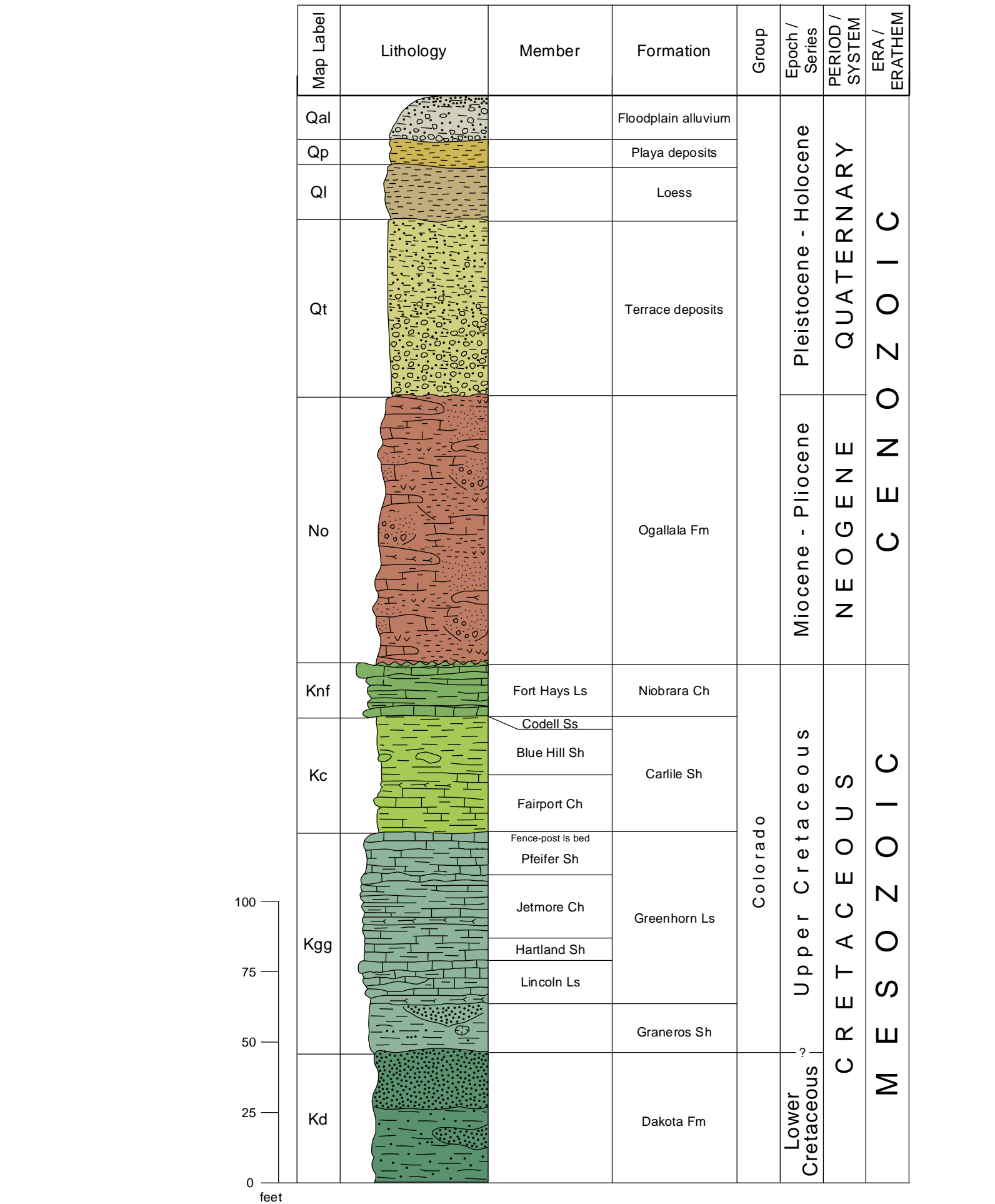
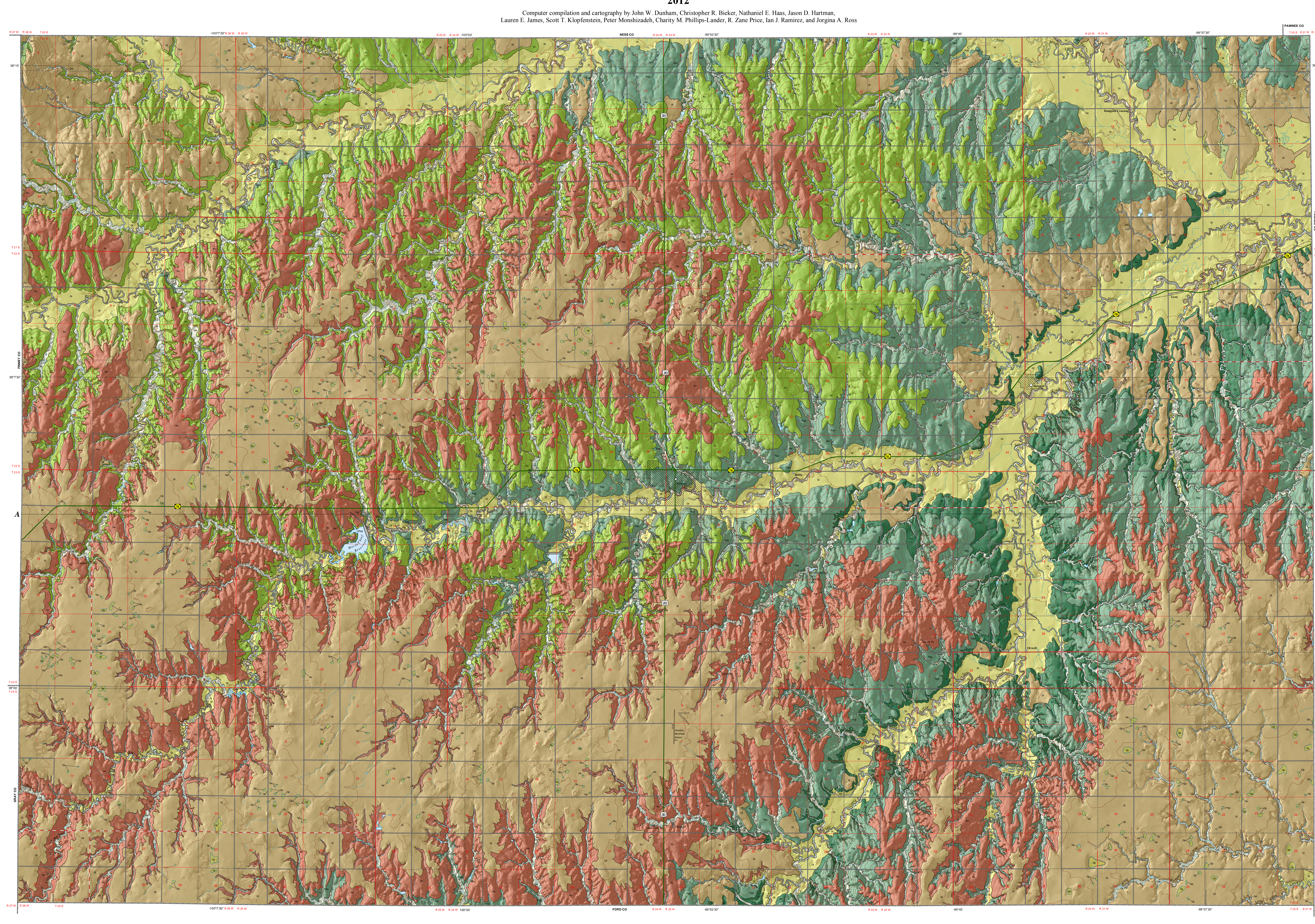
Oil and gas production is concentrated in subsurface Paleozoic rock formations located in all but the southwestern third of Hodgeman County. Over 45 million barrels of oil and 3 million net (1000 cubic ft) of gas have been produced in Hodgeman County. Oil production was in steady decline (601,660 barrels from 280 wells in 1995 to 303,644 barrels from 216 wells in 2007) until production started to increase in 2008. In 2011, 477,077 barrels of oil from 257 wells were reported. Gas production in the county has been intermittent, the best year since 1995 was 2007 when 584,900 net were produced from 19 wells. Since 2010, gas production has not been recorded in Hodgeman County (Kansas Geological Survey, 2012b).

WATER RESOURCES

Surface runoff of the major drainages in Hodgeman County is characterized by very low to no baseflow, except during significant precipitation events, and by high variability in yearly averages, ranging from years of no flow to years of high flow (Whitmore and Macfarlane, 2001; U.S. Geological Survey, 2012). Water for residential and agricultural uses comes from wells tapping ground water in the alluvium of the Pawnee River, Backster Creek, and Saw Log Creek, in the Ogallala Formation, and from sandstone in the Dakota Formation (Moss, 1932; Fabel, 1952; Kime and Spinnaker, 1985). The primary sources of water for agricultural irrigation are the Dakota and High Plains (Ogallala) aquifers; however, only about 9% of the total acres in the county are permitted for irrigation (Kansas Geological Survey, 2012b). The rest of the county has low potential for irrigation. It is not important for ground-water recharge and wetland habitat, and have functioned as perennial or ephemeral water sources for thousands of years (Howen et al., 2009).

EXPLANATION

Map legend including: Boundaries and Locations, Transportation, Geologic Unit Boundaries, Hydrology and Topography, Resource Development, and CITED REFERENCES.

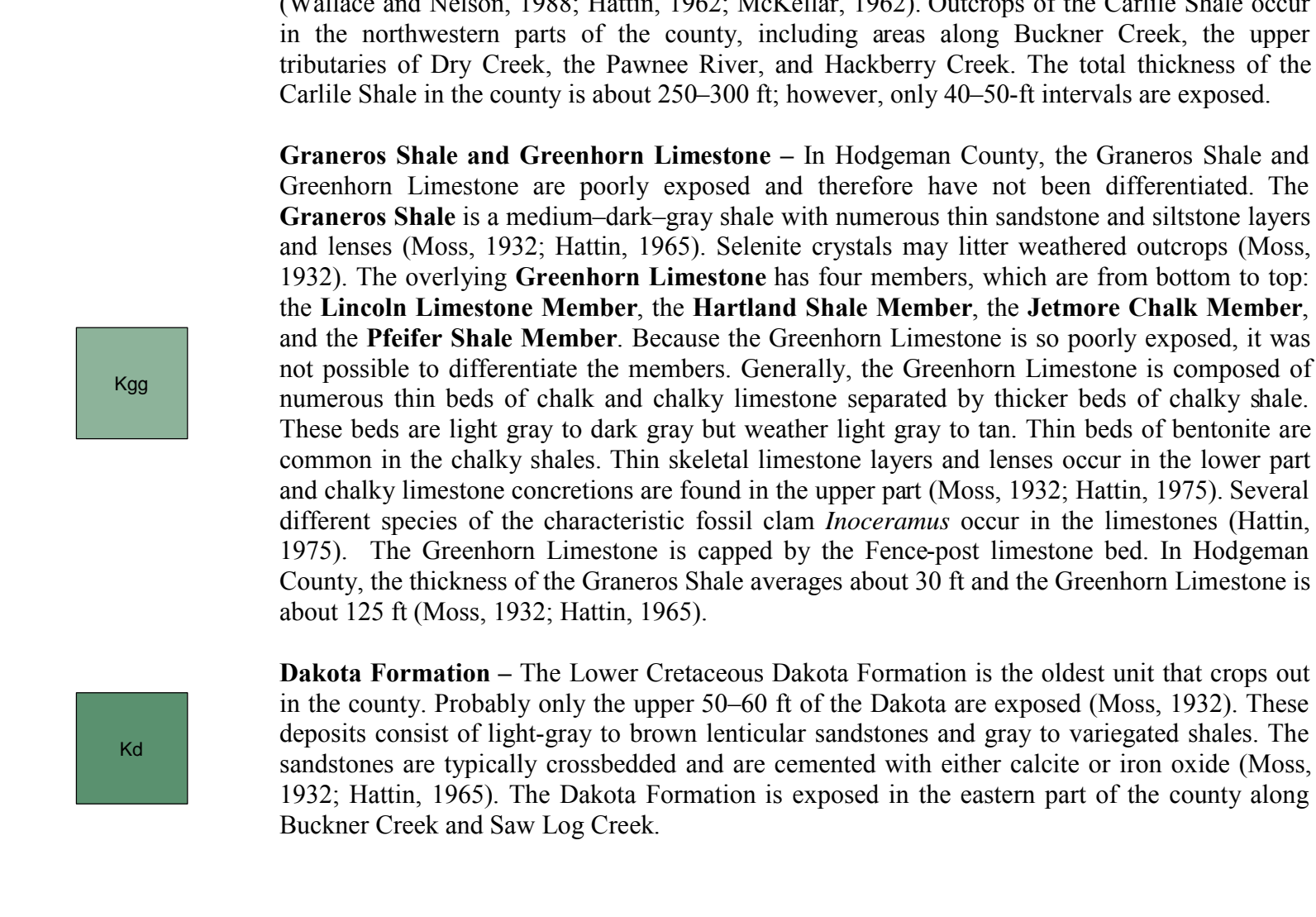


CENOZOIC ROCKS
Undifferentiated floodplain alluvium - Alluvial floodplain deposits are composed primarily of gravel and silty clay that were derived mostly from the surrounding Ogallala and Cretaceous bedrock. The average thickness of these deposits is about 14 ft.
Upland intermittent lake (playa) deposits - Shallow basins, also known as playas or buffalo wallows, have developed in the upland loess deposits, mainly south of the Pawnee River. The origin of these features is usually attributed to wind deflation, animal activity, dissolution, or some combination of these processes. The age of the playas probably ranges from late Pliocene to Holocene. The basins range in size from less than an acre to tens of acres. The basin fill consists of re-deposited silt and fine sand from the loess. In the larger basins, a caliche layer typically develops a few feet below the basin floor. The average thickness of the fill is approximately 6 ft.
Loess - Uplands in the county are mantled by loess composed of wind-deposited silt and minor amounts of clay and fine sand. The loess is calcareous and has a yellow color. The age of the loess ranges from late Pliocene to Holocene, and the thickness is up to 18 ft.
Alluvial terraces - Terraces occur along the Pawnee River and its major tributaries: Hackberry Creek, Backster Creek, and Saw Log Creek. Silt and clay dominate the fill, but the deposits also contain sand and gravel likely derived from the Ogallala Formation. The age of the terraces fill ranges from late Pliocene to at least middle Holocene. The average thickness of the terrace fill is about 60 ft.
Ogallala Formation - The Ogallala Formation is primarily Miocene and earliest Pliocene in age (Ludvigson et al., 2009) and is composed of gravel, sand, silt, and clay deposited by streams that transported sediments eastward from the Rocky Mountains. These sediments are variably cemented (ranging from unconformably to calcite-type deposits) with calcium carbonate. Throughout the Ogallala, thick calciche beds, referred to regionally as "mortar beds," regularly and discontinuously crop out (Moss, 1932). Silt dominates the Ogallala and commonly occurs in poorly sorted lenticular bodies. In Hodgeman County, the Ogallala Formation caps the uplands, is exposed along rivers and streams, and has a relatively uniform eastward dip of about 10 ft per mile (Moss, 1932). The maximum thickness of the Ogallala in the county is about 100 ft.

MESOZOIC ROCKS
Niobrara Chalk - The Upper Cretaceous Niobrara Chalk is composed of two members, the Smoky Hill Chalk Member and the underlying Fort Hays Limestone Member. As a result of pro-Ogallala erosion, the Smoky Hill Chalk Member has been completely eroded and only the Fort Hays Limestone Member is exposed in Hodgeman County (Moss, 1932). The Fort Hays Limestone Member is composed of thick beds of light-to-dark-gray chalk separated by thin beds of cherty shale. Weathered outcrops are cream-to-tan colored. The only occurrence of Fort Hays limestone in the extreme northwest corner of the county is in an area less than half of a square mile (Moss, 1932). Only about 20 ft of Fort Hays limestone is exposed in Hodgeman County.
Carlisle Shale - The Cretaceous Carlisle Shale is composed of three members, in ascending order: the Fairport Chalk Member, the Blue Hill Shale Member, and the Coddell Sandstone Member. The Fairport Chalk Member is an olive-gray to dark-gray, blocky, fossiliferous, cherty shale, intercalated with cherty and cherty limestone beds and a few thin bentonite beds. Thin, fossiliferous cherty limestone are more common near the base of the Fairport chalk and form small benches when weathered (Moss, 1932; Hattin, 1962). The upper part of the Carlisle Shale is represented by the Blue Hill Shale Member, which is a blue-shale, blocky to fissile, slightly silty shale containing pyrite nodules, selenite crystals, and septarian concretions (Moss, 1932; Hattin, 1962). The contact between the Blue Hill shale and Fairport chalk is not exposed in Hodgeman County (Moss, 1932). The Coddell Sandstone Member has not been observed in the county (Wallace and Nelson, 1988; Hattin, 1962; McKellar, 1962). Outcrops of the Carlisle Shale occur in the northwestern parts of the county, including areas along Backster Creek, the upper tributaries of Dry Creek, the Pawnee River, and Hackberry Creek. The total thickness of the Carlisle Shale in the county is about 250 ft; however, only 50-ft intervals are exposed.

Greenhorn Shale and Greenhorn Limestone - In Hodgeman County, the Greenhorn Shale and Greenhorn Limestone are poorly exposed and therefore have not been differentiated. The Greenhorn Shale is a medium-dark gray shale with numerous thin sandstone and siltstone layers and lenses (Moss, 1932; Hattin, 1962). Selenite crystals may form weathered outcrops (Moss, 1932). The overlying Greenhorn Limestone has four members, which are from bottom to top: the Lincoln Limestone Member, the Hartland Shale Member, the Armand Chalk Member, and the Pfeiffer Shale Member. Because the Greenhorn Limestone is so poorly exposed, it was not possible to differentiate the members. Generally, the Greenhorn Limestone is composed of numerous thin beds of chalk and cherty limestone, separated by thicker beds of cherty shale. These beds are light gray to dark gray but weather light gray to tan. Thin beds of bentonite are common in the cherty shales. Thin selenite concretions and lenses occur in the lower part and cherty limestone concretions are found in the upper part (Moss, 1932; Hattin, 1975). Several different species of the characteristic fossil clam Anserinus occur in the limestones (Hattin, 1975). The Greenhorn Limestone is capped by the Forcepoint limestone bed. In Hodgeman County, the thickness of the Greenhorn Shale averages about 30 ft and the Greenhorn Limestone is about 12.5 ft (Moss, 1932; Hattin, 1962).

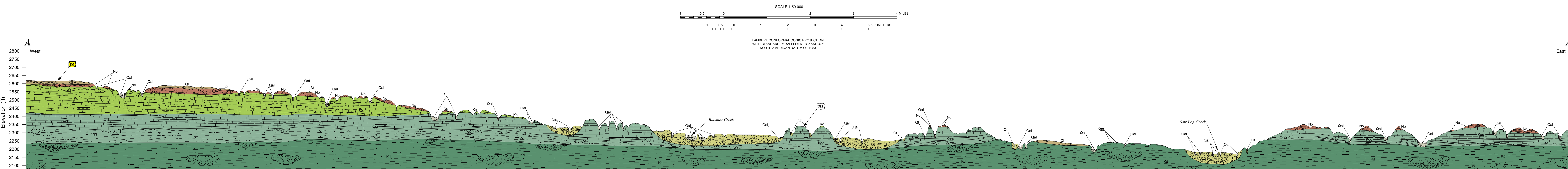
Dakota Formation - The Lower Cretaceous Dakota Formation is the oldest unit that crops out in the county. Probably only the upper 50-60 ft of the Dakota are exposed (Moss, 1932). These deposits consist of light-gray to brown lenticular sandstones and gray to variegated shales. The sandstones are typically crossbedded and are cemented with either calcite or iron oxide (Moss, 1932; Hattin, 1962). The Dakota Formation is exposed in the eastern part of the county along Backster Creek and Saw Log Creek.



Scale bars: 1:50,000 and 1:100,000. Vertical exaggeration: 20x.

ADDITIONAL SOURCES
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SUGGESTED REFERENCE TO THIS MAP
Johnson, W. C., and Woodburn, T. L., 2012. Surficial geology of Hodgeman County, Kansas. Kansas Geological Survey, Map M-107, scale 1:50,000.



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Publication Series
(F80) 864-2157
or visit the Kansas Geological Survey website at www.kgs.ku.edu
ISBN # 1-55886-941-8
Last updated 11/28/2012